

Interpretation of Gardner's equation

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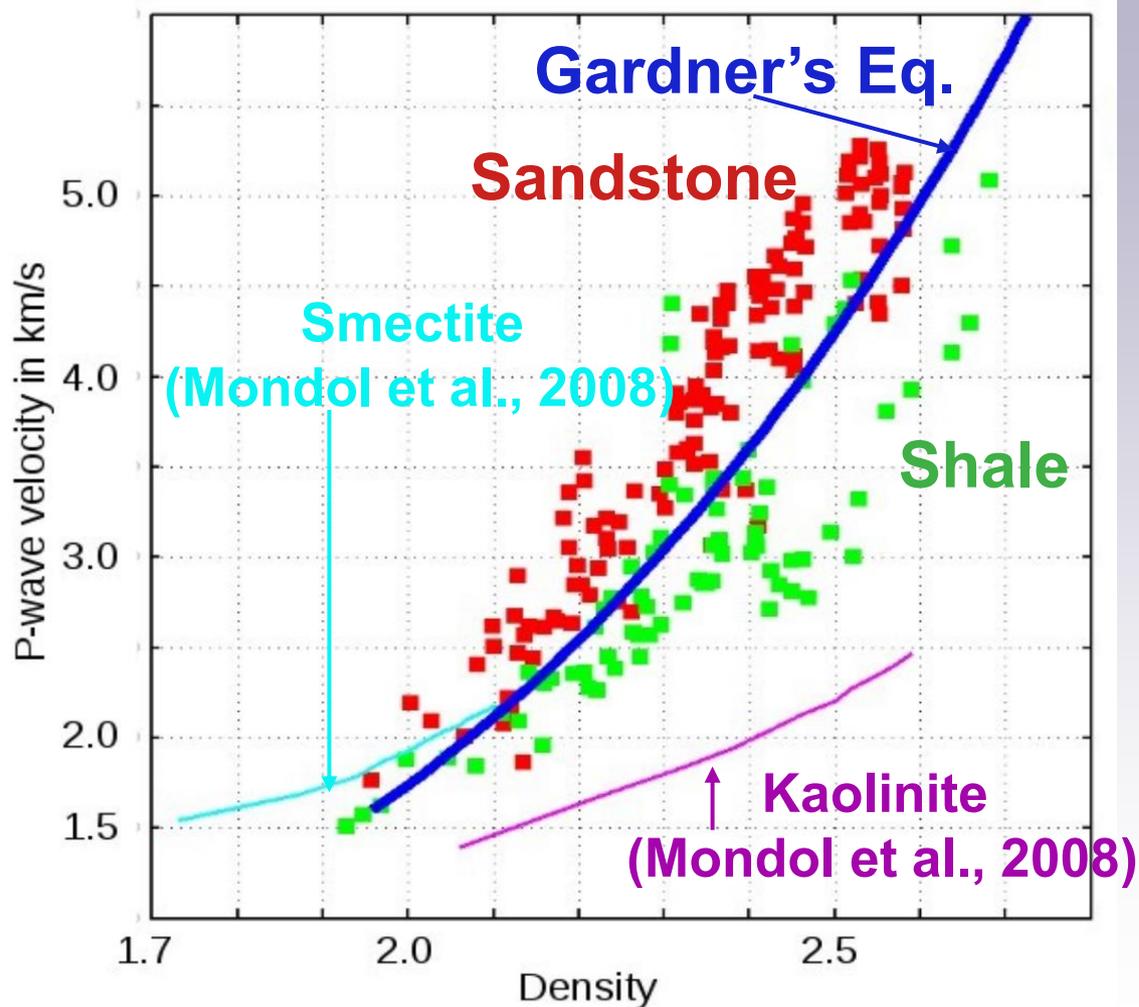
Defined Quantities

- V_p P-wave velocity
- V_b Brine P-wave velocity (~1500 m/s)
- V_s S-wave velocity
- ρ Density
- ϕ Porosity
- K Bulk Modulus
- μ Shear Modulus

Subscripts

- w wet – fluid filled pores
- d dry - no fluid in pores
- m mineral
- f fluid
- c critical

Gardner's Equation



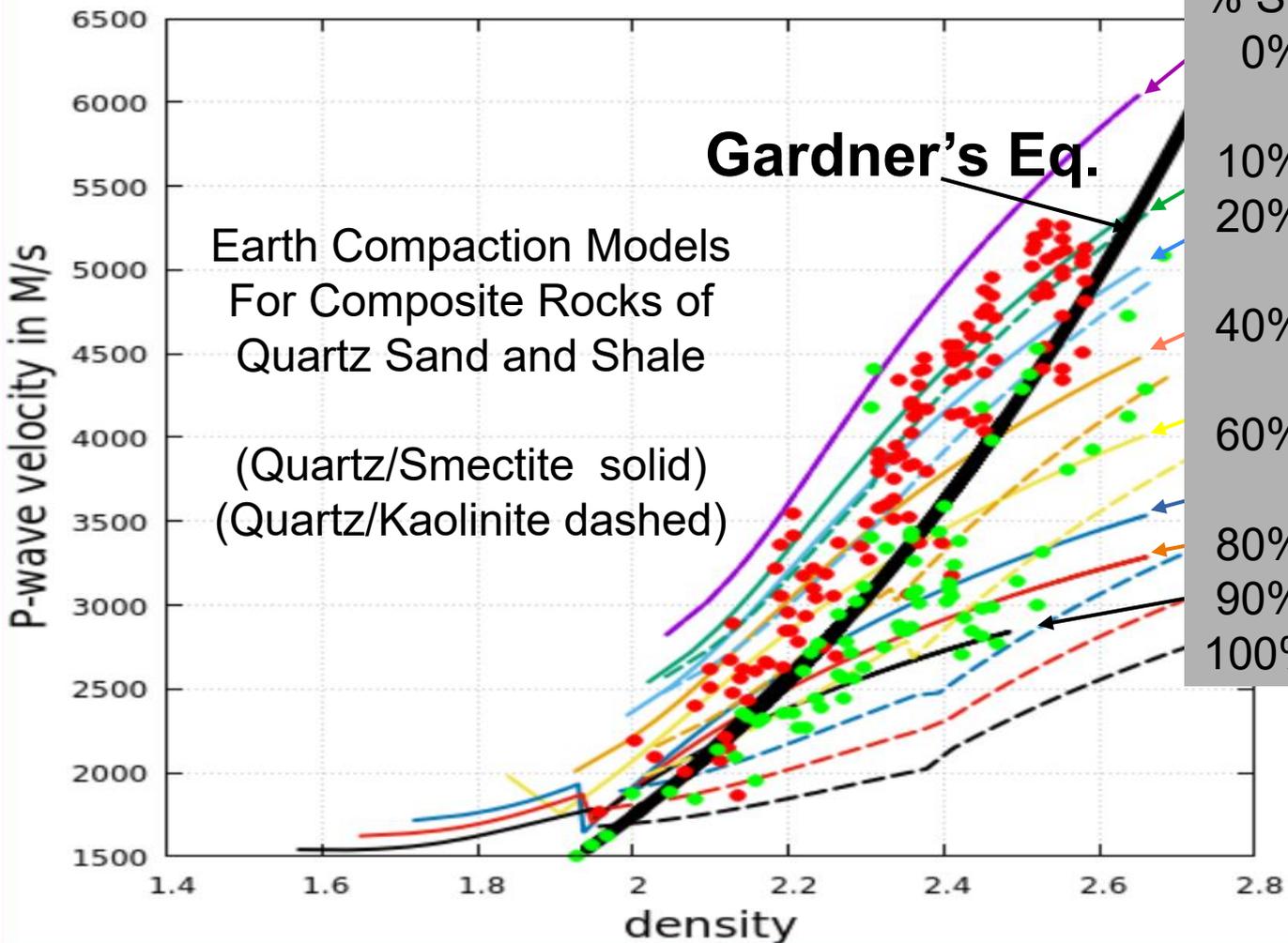
Gardner's Equation (Gardner et al., 1974, equation 7):

$$\rho = \rho_0 V^{\frac{1}{4}}, \quad \rho_0 = (1.741 \text{ Km/s})^{-1/4}$$

Density in g/cc, Sandstone and Shale data points from Castagna et al. (1993, Fig 9).

The fit is not to the points shown but instead It is a fit to "... the more prevalent sedimentary rock types through a wide range of basins, geologic ages, and depths (to 25,000 ft)"

P-wave velocity versus Density



% S

0%
10%
20%
40%
60%
80%
90%
100%

% K

0%
10%
20%
40%
60%
80%
90%
100%

Smectite

Kaolinite

Compaction Modeling

Used To Analyze Gardner's Eq.

- Elastic Wave Propagation definitions of V_p and V_s
 - Nur's Critical Porosity model
 - Gassmann's Equations
 - Aplin's Equation relating porosity to effective stress
-
- Density vs velocity from earth compaction modeling by
combining Nur with Aplin
 - A different density vs velocity relation
that results from *combining Nur with Gassman*

Combining Gassmann's Equations with Nur's Critical Porosity Model Leads To These Equations

Mavko and Mukerji (1998), Krief (1990),
Higginbotham et al. (2011, 2012)

$$\frac{V_p^2}{a^2} - \frac{V_s^2}{b^2} = 1$$

a, b, s, C_ρ

Higginbotham et al. (2012)

$$\rho = \frac{C_\rho}{1 - \left(\frac{sV_p}{V_b}\right)^2}$$

$$V_b = 1500 \text{ m/s}$$

all constant for a given mineral and pore fluid.

Combining Gassmann's Equations with Nur's Critical Porosity Model Leads To These Equations

Mavko and Mukerji (1998), Krief (1990),
Higginbotham et al. (2011, 2012)

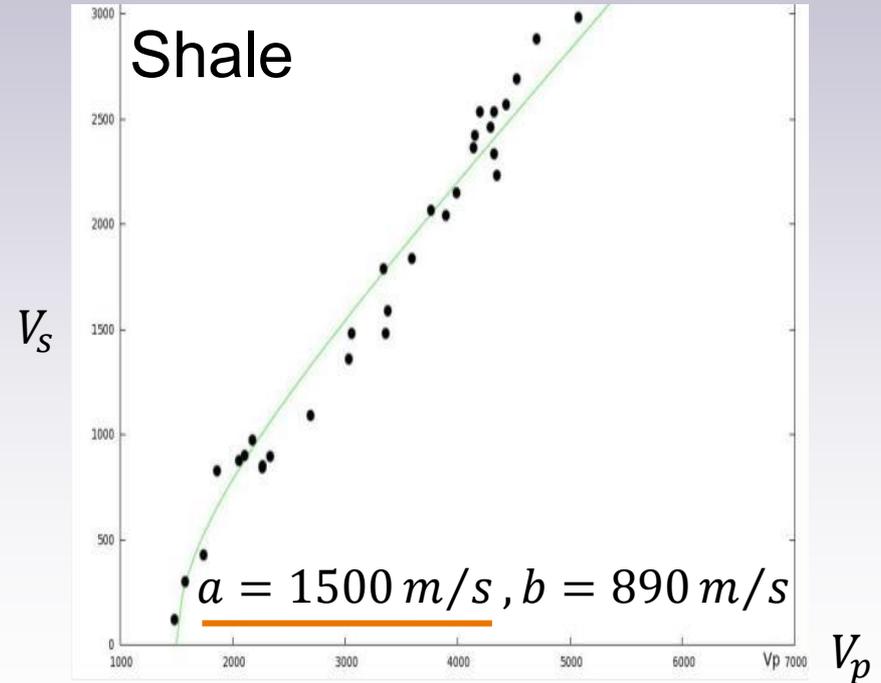
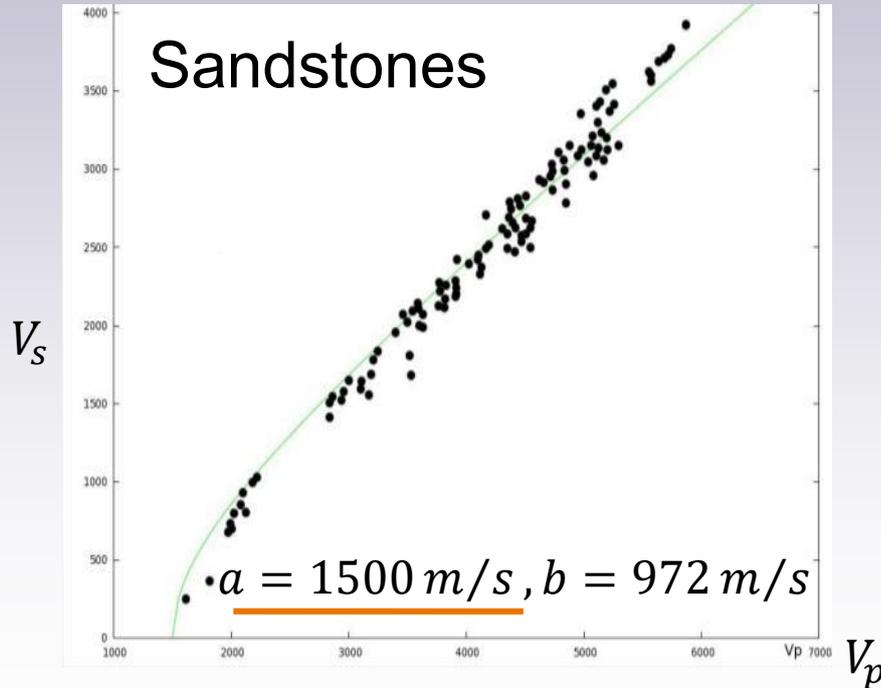
$$\frac{V_p^2}{a^2} - \frac{V_s^2}{b^2} = 1$$

Brief attention for
this equation

EVIDENCE

$$\frac{V_p^2}{a^2} - \frac{V_s^2}{b^2} = 1$$

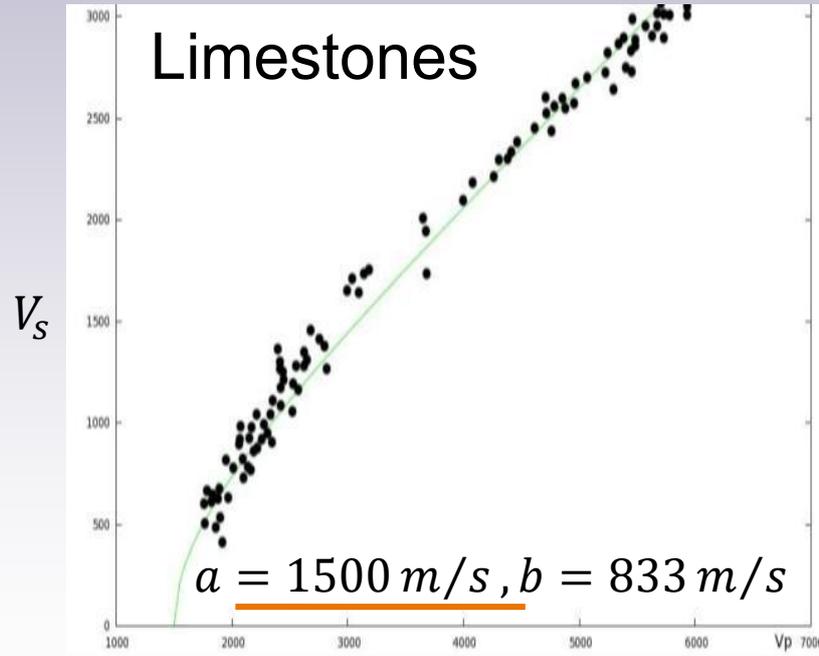
These plots are from Higginbotham (2011), data points from Castagna (1993, p 138-139).



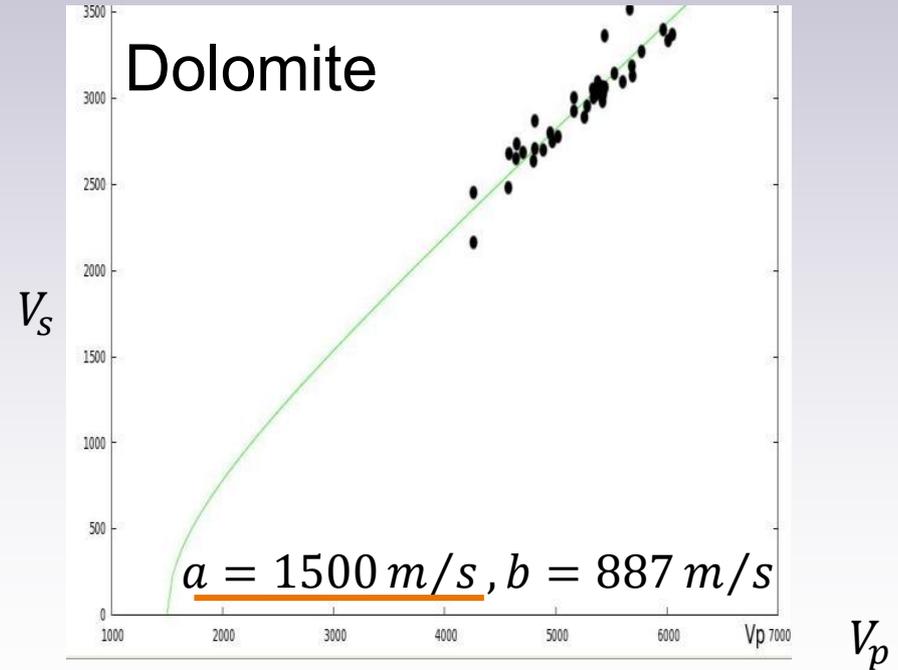
EVIDENCE

$$\frac{V_p^2}{a^2} - \frac{V_s^2}{b^2} = 1$$

These plots are from Higginbotham (2011), data points from Castagna (1993, p 138-139).



V_p

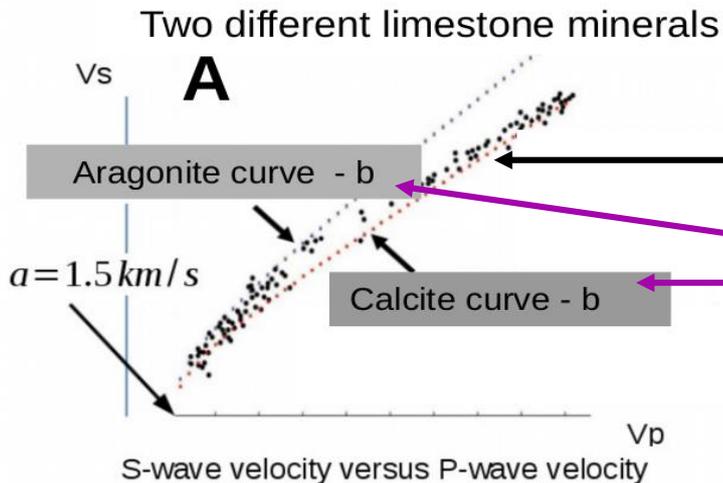


V_p

EVIDENCE

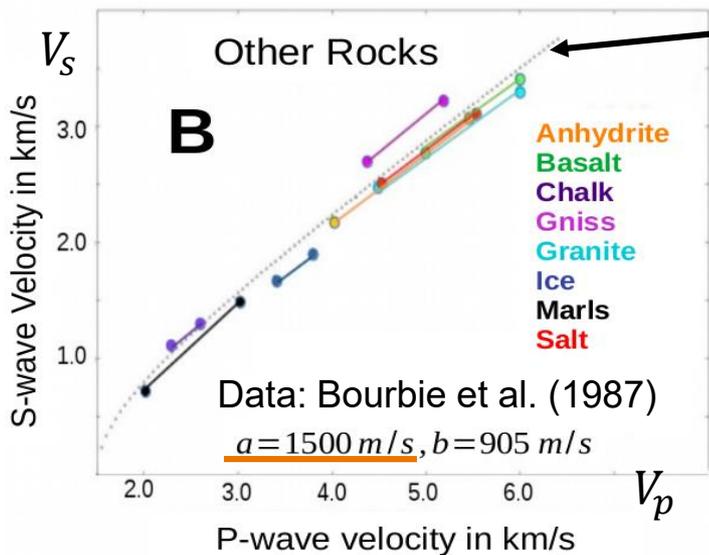
continued

Scatter is probably due to impurities.



$$\frac{1}{b^2} = \frac{1}{a^2} \left(\frac{4}{3} + \frac{K_m}{\mu_m} \right) - \frac{\rho_m}{\mu_m}$$

Note: Faint dashed curve is a fit for $b = 905 \text{ m/s}$ to all the data shown in the previous slide – not a fit to the data points plotted in Figure B.



Appears to work for a variety of rocks even with a fixed at 1500 m/s.

Gassmann's Eq's are not supposed to work for these rocks ???

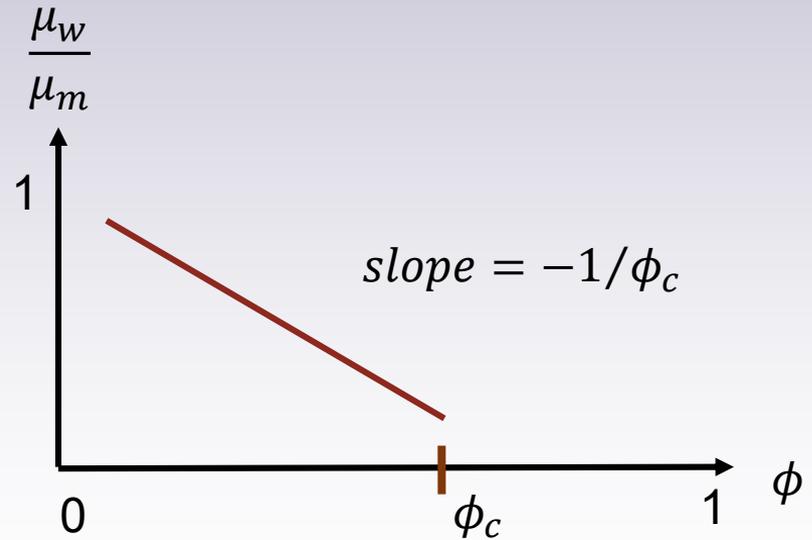
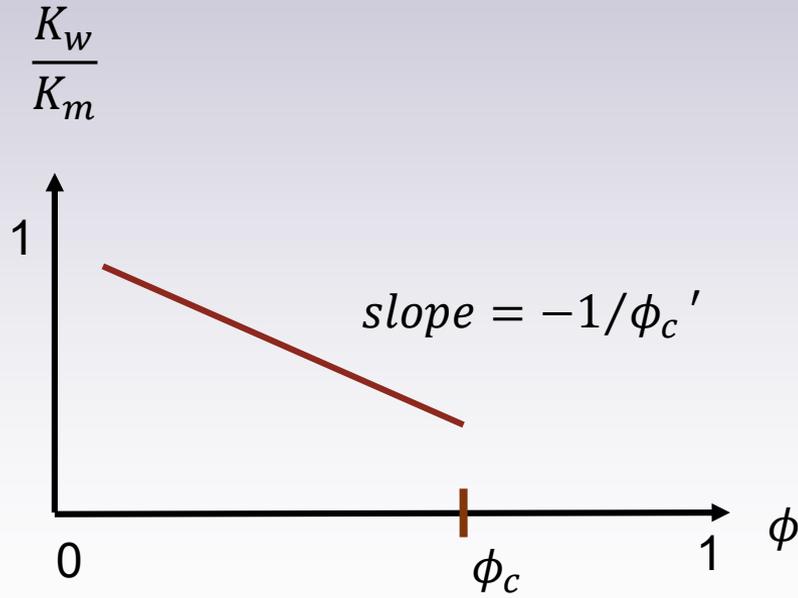
Quote from Mavko et al. (2009, pg 273)

*Gassmann's equation ... is valid only at sufficiently **low frequencies** ... such that there is sufficient time for the pore fluid to flow and eliminate wave-induced pore pressure gradients ...*

Combining Gassmann's Equations with Nur's Critical Porosity Model

$$K_w/K_m = \left(1 + \left(\frac{-1}{\phi_c'} \right) \phi \right)$$

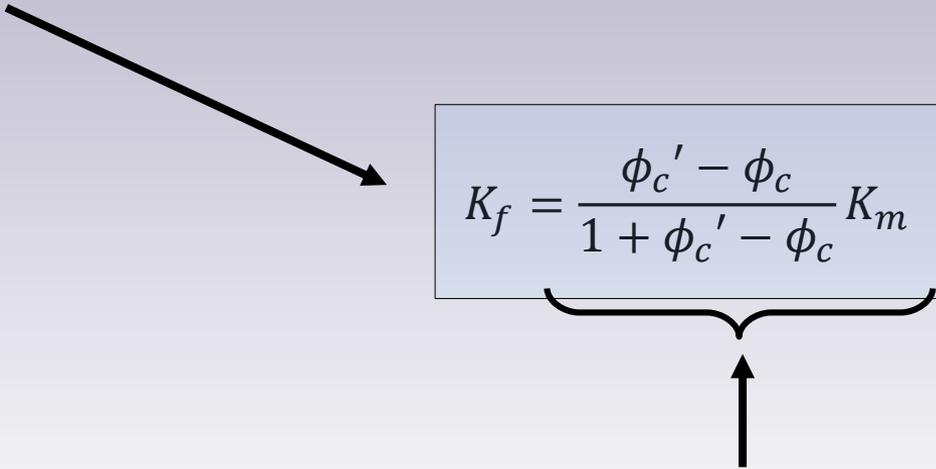
$$\mu_w/\mu_m = \left(1 + \left(\frac{-1}{\phi_c} \right) \phi \right)$$



Gassman-Nur also provides

$$\phi_c' = \phi_c + \frac{K_f}{K_m - K_f}$$

Higginbotham et al. (2012)


$$K_f = \frac{\phi_c' - \phi_c}{1 + \phi_c' - \phi_c} K_m$$

The **MODIFIED** Gassman-Nur Model replaces K_f with this expression.

THE FLUID BULK MODULUS VANISHES FROM THE EQUATIONS

It is reasonable to conclude that the fluid bulk modulus K_f is the source of error in Gassmann's Eqs for low permeable rocks and high frequencies.

Eliminate K_f

$$K_f \rightarrow \frac{\phi_c' - \phi_c}{1 + \phi_c' - \phi_c} K_m$$

Removing Dependence on Fluid Bulk Modulus

$$K_f \rightarrow \frac{\phi_c' - \phi_c}{1 + \phi_c' - \phi_c} K_m$$

Modified Gassman-Nur Model
No dependence on K_f

$$a = a(\phi_c, \phi_c', K_f, K_m, \rho_f, \rho_m)$$

→

$$a = a(\phi_c, \phi_c', K_m, \rho_f, \rho_m)$$

$$b = b(\phi_c, \phi_c', K_f, K_m, \rho_f, \rho_m, \mu_m)$$

→

$$b = b(\phi_c, \phi_c', K_m, \rho_f, \rho_m, \mu_m)$$

$$C_\rho = C_\rho(\phi_c, \phi_c', K_f, K_m, \rho_f, \rho_m, \mu_m)$$

→

$$C_\rho = C_\rho(\phi_c, \phi_c', K_m, \rho_f, \rho_m, \mu_m)$$

$$s = s(\phi_c, \phi_c', K_f, K_m, \rho_f, \rho_m, \mu_m)$$

→

$$s = s(\phi_c, \phi_c', K_m, \rho_f, \rho_m, \mu_m)$$

Now Focus on the Density Eq.

$$\frac{V_p^2}{a^2} - \frac{V_s^2}{b^2} = 1$$

$$\rho = \frac{C_\rho}{1 - \left(\frac{sV_p}{V_b}\right)^2}$$

$$V_b = 1500 \text{ m/s}$$

Smectite

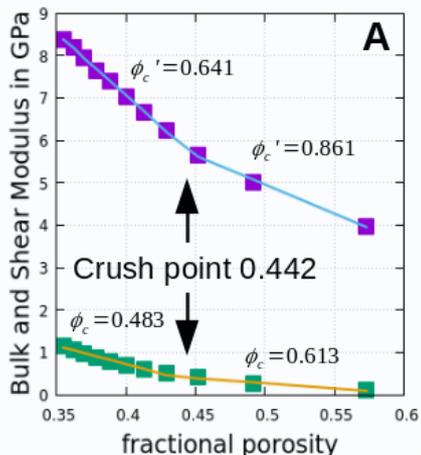
Kaolinite

Shale

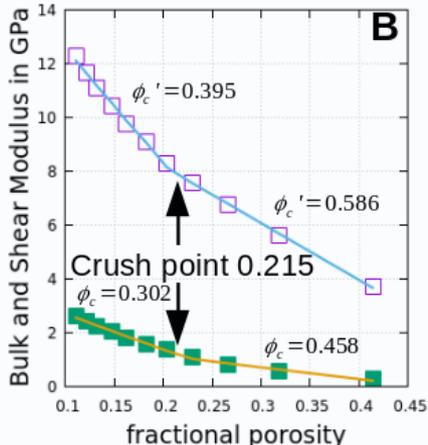
Ultrasonic Freq.

- Mondol's data (Mondol et al., 2008) can be fit quite well to *two connected straight lines*.
- The crush point (grain reorientation) may be the consolidation threshold porosity of Vernik et al. (2010a, 2010b).
- Elastic properties change suddenly at this crush point (consolidation threshold).
- **Nur's model holds** on either side of the crush point porosity!
- Bulk modulus slope supplies ϕ_c' (slide 12)
- Shear modulus slope supplies ϕ_c (slide 12)

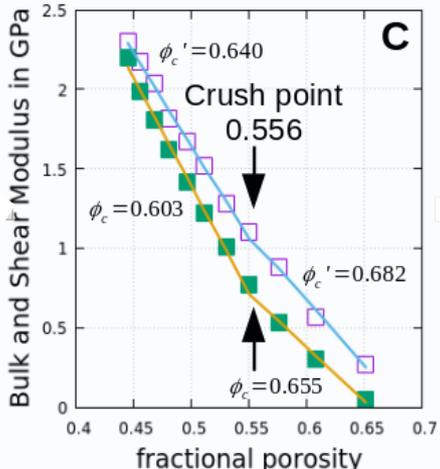
Smectite - Brine saturated



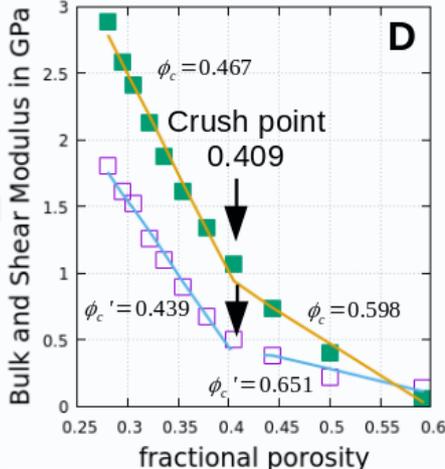
Kaolinite - Brine saturated



Smectite - dry



Kaolinite - dry



Evidence SMECTITE

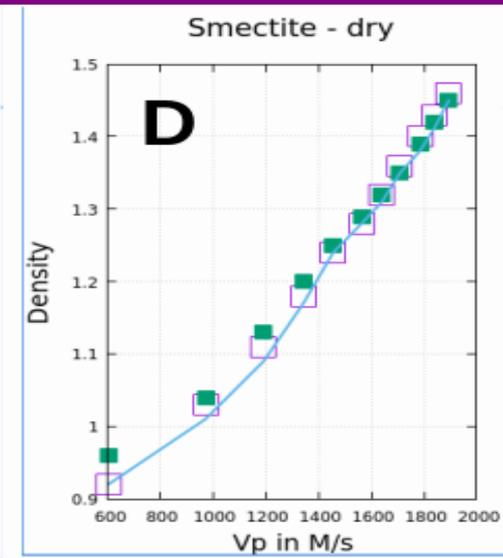
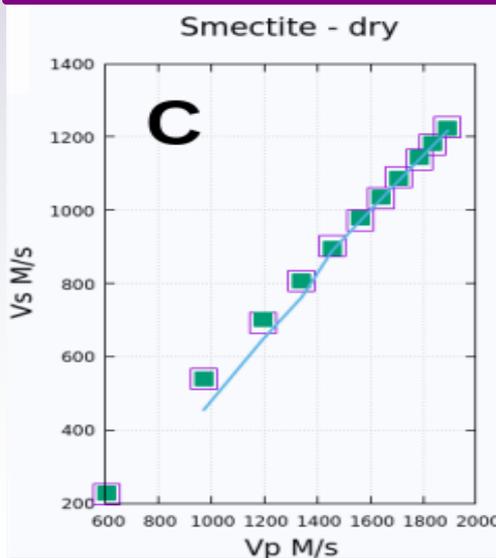
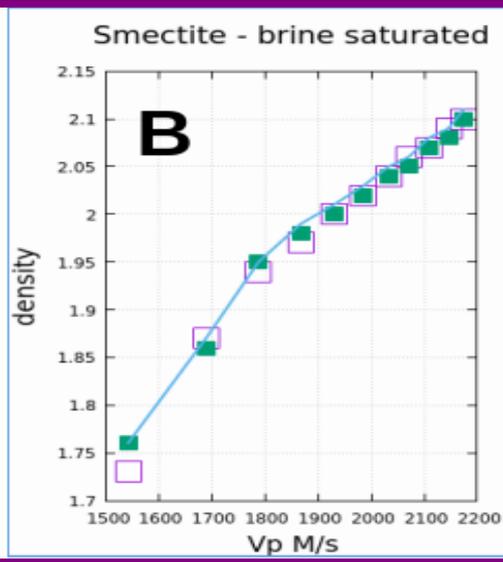
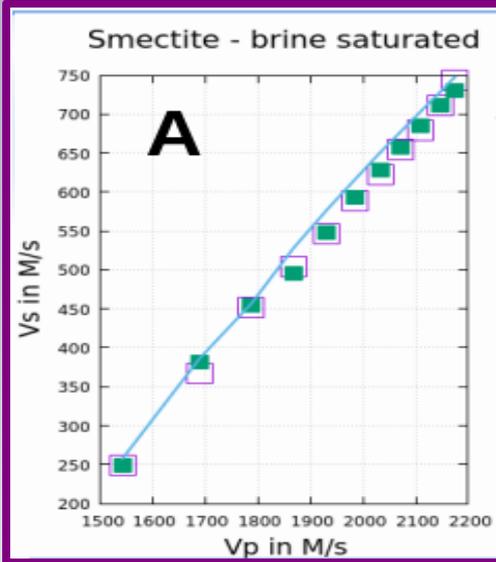
Open squares measured points
(Mondol et al., 2008)

Green: Modified Gassman-Nur

Blue: Gassman-Nur using K_f

Wet grain density increased by 2%
for 89% pure sample.

Dry grain density decreased by
3% for 89% pure sample.



Evidence Kaolinite

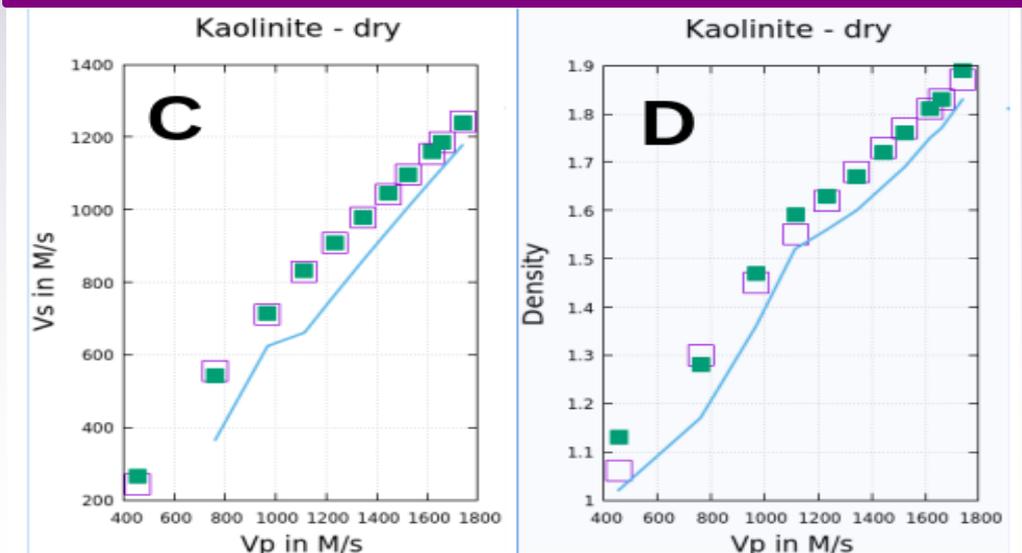
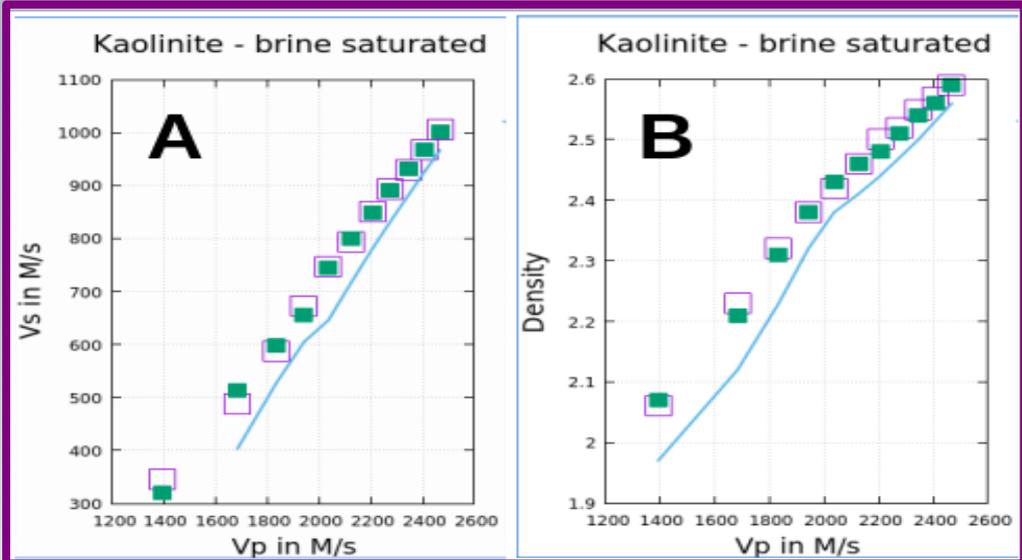
Open squares measured points
(Mondol et al., 2008)

Green: Modified Gassman-Nur

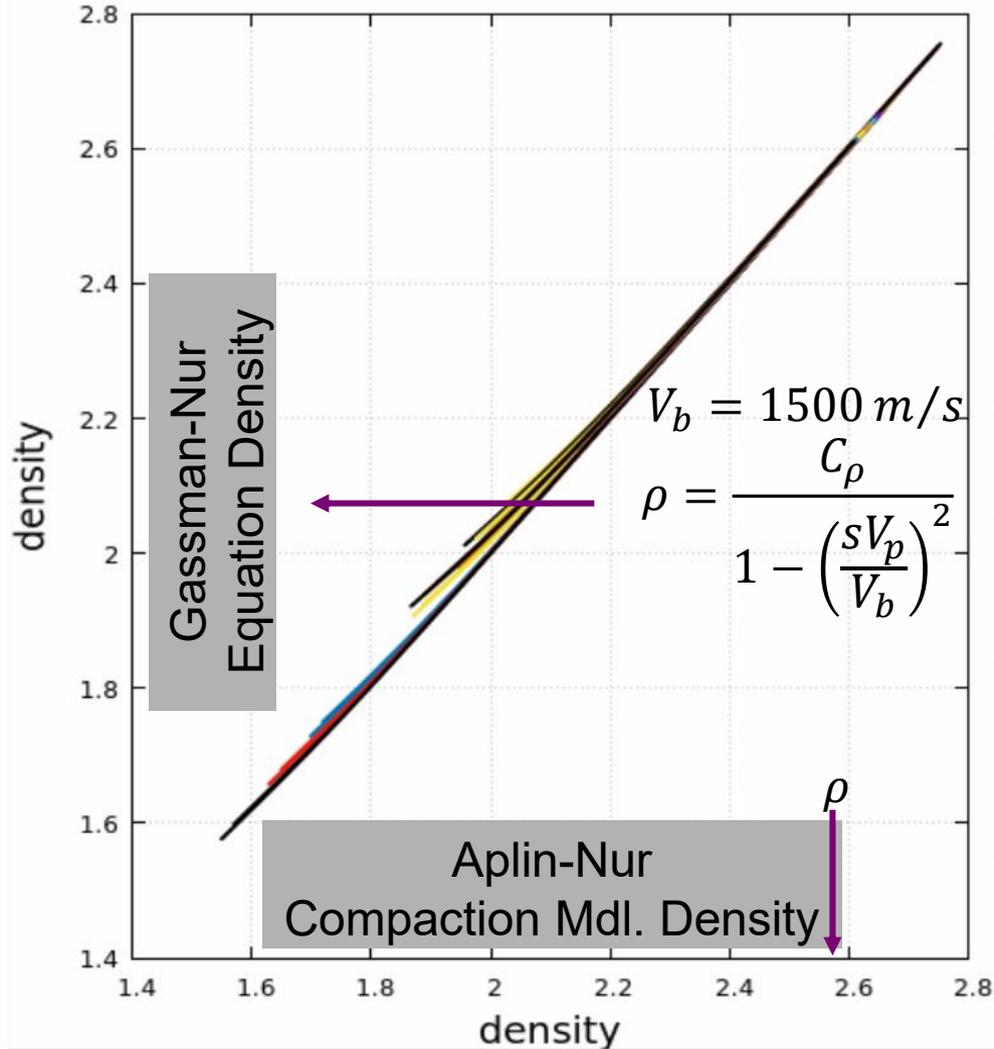
Blue: Gassman-Nur using K_f

Wet grain density increased by 6%
for 81% pure sample.

Dry grain density decreased by
2% for 81% pure sample.



Computed Density versus Density



- 0 S, 0.08, E
- 10 S, 0.09, E
- 20 S, 0.10, E
- 30 S, 0.10, E
- 50 S, 0.13, E
- 80 S, 0.14, E
- 90 S, 0.15, E
- 100 S, 0.16, E
- 0 K, 0.08, E
- 10 K, 0.09, E
- 20 K, 0.10, E
- 30 K, 0.10, E
- 50 K, 0.13, E
- 80 K, 0.14, E
- 90 K, 0.15, E
- 100 K, 0.16, E
- 0 S, 0.08
- 10 S, 0.09
- 20 S, 0.10
- 30 S, 0.10
- 50 S, 0.13
- 80 S, 0.14
- 90 S, 0.15
- 100 S, 0.16
- 0 K, 0.08
- 10 K, 0.09
- 20 K, 0.10
- 30 K, 0.10
- 50 K, 0.13
- 80 K, 0.14
- 90 K, 0.15
- 100 K, 0.16

S.

K.

S.

K.

Modified Gassmann-Nur

Gassmann-Nur

Conclusions

- The earth compaction model agrees with Gardner's equation for composite sand shale rocks having high percentage sand content
- The earth compaction model does not agree with Gardner's equation for composite sand shale rocks having high percentage shale content.
- Evidence is presented that supports the use of equations derived from Gassmann's equations combined with Nur's critical porosity model for high frequency low permeability rocks.
- The equation relating density to P-wave velocity derived from combining Gassmann's equations with Nur's critical porosity model is shown to agree closely with the compaction models based on combining Nur's model with Aplin's equation.
- *Note: The last two slides provide the formulas used for Bulk Modulus and Density in the compaction models as a function of temperature, salinity, and pressure.*

Bulk Modulus

Bulk modulus of brine in bars as a Fortran function of temperature, $20 < T < 350$, in degrees C, salinity, $1 < S < 240,000$, in parts per million, and pressure, $98.1 < P < 981$, in bars, from a fit to Castagna et al. (1993, Fig 21):

$$\begin{aligned} \text{BlkMod_of_T_S_P}(T,S,P) = & \\ & : 0.1924\text{E}+05 + 0.6646\text{E}+01 * P - 0.6039\text{E}-03 * P^{**2} + 0.1650\text{E}+03 * T \\ & : -0.3031\text{E}-01 * T * P + 0.1759\text{E}-04 * T * P^{**2} - 0.1671\text{E}+01 * T^{**2} \\ & : + 0.5463\text{E}-03 * T^{**2} * P - 0.2344\text{E}-06 * T^{**2} * P^{**2} + 0.4003\text{E}-02 * T^{**3} \\ & : -0.4519\text{E}-06 * T^{**3} * P - 0.3545\text{E}-05 * T^{**4} + 0.4100\text{E}-01 * S \\ & : + 0.1776\text{E}-04 * S * P - 0.4226\text{E}-08 * S * P^{**2} - 0.2771\text{E}-03 * T * S \\ & : -0.6566\text{E}-07 * T * S * P + 0.1577\text{E}-10 * T * S * P^{**2} + 0.1476\text{E}-05 * T^{**2} * S \\ & : + 0.1796\text{E}-09 * T^{**2} * S * P - 0.2970\text{E}-08 * T^{**3} * S + 0.1031\text{E}-06 * S^{**2} \\ & : -0.2194\text{E}-10 * S^{**2} * P + 0.1324\text{E}-13 * S^{**2} * P^{**2} - 0.1425\text{E}-09 * T * S^{**2} \\ & : -0.7629\text{E}-15 * T * S^{**2} * P + 0.2335\text{E}-12 * T^{**2} * S^{**2} . \end{aligned}$$

Density

Density of brine in bars as a Fortran function of temperature, $20 < T < 350$, in degrees C, salinity, $1 < S < 300,000$, in parts per million, and pressure, $1 < P < 1000$, in bars, from a fit to Castagna et al. (1993, Fig 20):

$$\begin{aligned} \text{Brine_D_of_T_S_P}(T,S,P) = & 0.21274\text{E}+02 * S^{**}(-2) \\ & :+ 0.16674\text{E}-01 * S^{**}(-2) * P - 0.15228\text{E}-04 * S^{**}(-2) * P^{**}2 \\ & :+ 0.10316\text{E}+00 * T * S^{**}(-2) - 0.40792\text{E}-03 * T * S^{**}(-2) * P \\ & :+ 0.28547\text{E}-06 * T * S^{**}(-2) * P^{**}2 - 0.77854\text{E}-03 * T^{**}2 * S^{**}(-2) \\ & :+ 0.20027\text{E}-05 * T^{**}2 * S^{**}(-2) * P - 0.83005\text{E}-09 * T^{**}2 * S^{**}(-2) * P^{**}2 \\ & :+ 0.19105\text{E}-05 * T^{**}3 * S^{**}(-2) - 0.27537\text{E}-08 * T^{**}3 * S^{**}(-2) * P \\ & :+ 0.11955\text{E}-13 * T^{**}3 * S^{**}(-2) * P^{**}2 - 0.21329\text{E}+02 * S^{**}(-1) \\ & :- 0.16670\text{E}-01 * S^{**}(-1) * P + 0.15228\text{E}-04 * S^{**}(-1) * P^{**}2 \\ & :- 0.10300\text{E}+00 * T * S^{**}(-1) + 0.40750\text{E}-03 * T * S^{**}(-1) * P \\ & :- 0.28515\text{E}-06 * T * S^{**}(-1) * P^{**}2 + 0.77699\text{E}-03 * T^{**}2 * S^{**}(-1) \\ & :- 0.19976\text{E}-05 * T^{**}2 * S^{**}(-1) * P + 0.82582\text{E}-09 * T^{**}2 * S^{**}(-1) * P^{**}2 \\ & :- 0.19067\text{E}-05 * T^{**}3 * S^{**}(-1) + 0.27394\text{E}-08 * T^{**}3 * S^{**}(-1) * P \\ & :+ 0.10624\text{E}+01 + 0.43787\text{E}-04 * P - 0.65442\text{E}-08 * P^{**}2 \\ & :- 0.40447\text{E}-03 * T - 0.57638\text{E}-07 * T * P - 0.27188\text{E}-11 * T * P^{**}2 \\ & :- 0.14361\text{E}-05 * T^{**}2 + 0.86305\text{E}-09 * T^{**}2 * P - 0.98617\text{E}-09 * T^{**}3 . \\ & :+ 0.52239\text{E}-06 * S + 0.35447\text{E}-10 * S * P \\ & :- 0.18153\text{E}-13 * S * P^{**}2 - 0.12925\text{E}-09 * T * S - 0.47644\text{E}-12 * T * S * P \\ & :+ 0.12919\text{E}-11 * T^{**}2 * S \end{aligned}$$